

Deformation and Strain

Deformation and Strain

Earth Structure (2019)
(Processes in Structural Geology & Tectonics)

© Ben van der Pluijm
2/15/2019 10:25

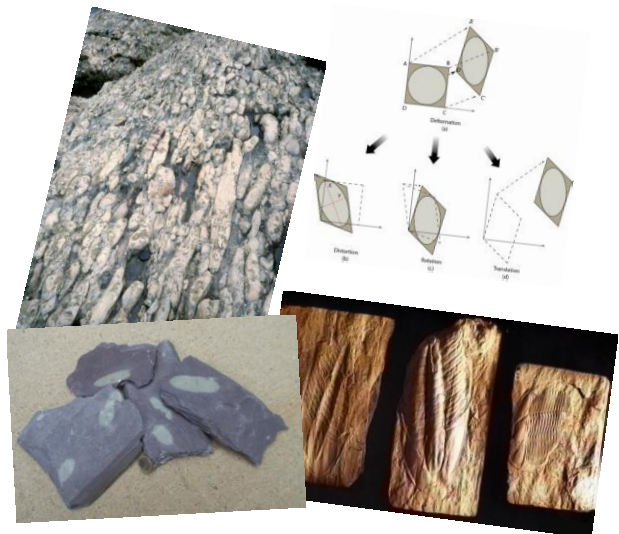
Earth Structure (2019)
(Processes in Structural Geology & Tectonics)

© Ben van der Pluijm
2/21/2019 16:32

We Discuss ...

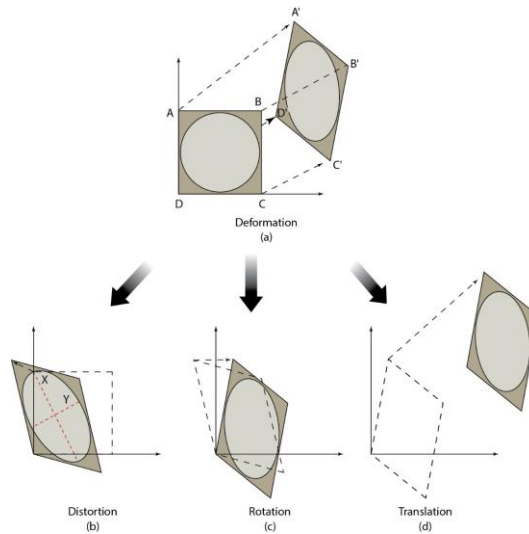
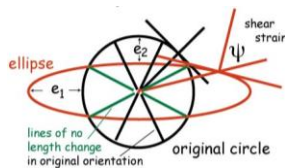
Deformation and Strain

- Deformation Components
- Homogeneous vs. Heterogeneous Strain
- Strain Ellipse (and Ellipsoid)
- Coaxial and Non-coaxial Strain
- Superposed Strain
- Strain Quantities
- Representations of Strain
- (Finite) Strain Analysis
 - Spherical Objects
 - Angular Changes
 - Length Changes
- Understanding Strain Values



The Components of Deformation

1. Strain (distortion)
 - a) Extension (or stretch)
length changes
 - b) Internal rotation (vorticity)
finite strain axes rotate relative to instantaneous strain axes
 - c) Volume change
2. Rigid-body rotation (or spin)
instantaneous strain axes and finite strain axes rotate together
3. Rigid-body translation



Homogeneous vs. Heterogeneous Strain

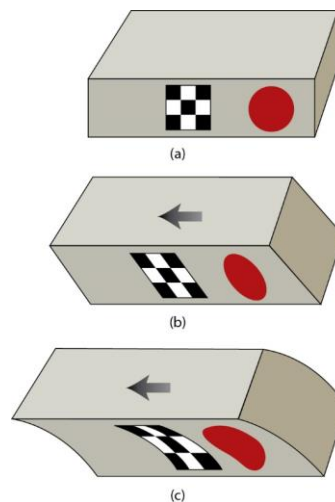
Homogeneous strain:

- Straight lines remain straight
- Parallel lines remain parallel
- Circles become ellipses (or spheres become ellipsoids)

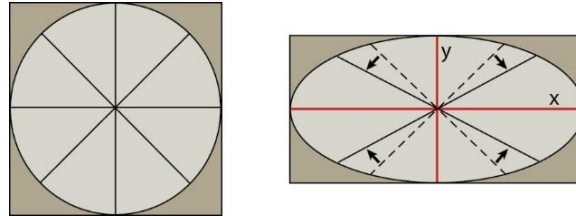
vs.

Heterogeneous strain

(deck of cards)

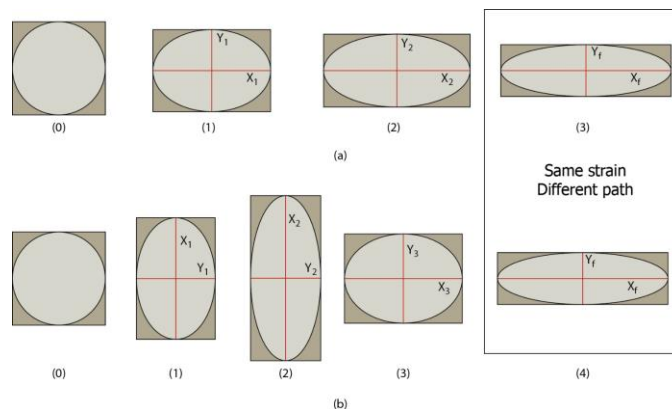


Homogeneous Strain: Principal Strain Axes



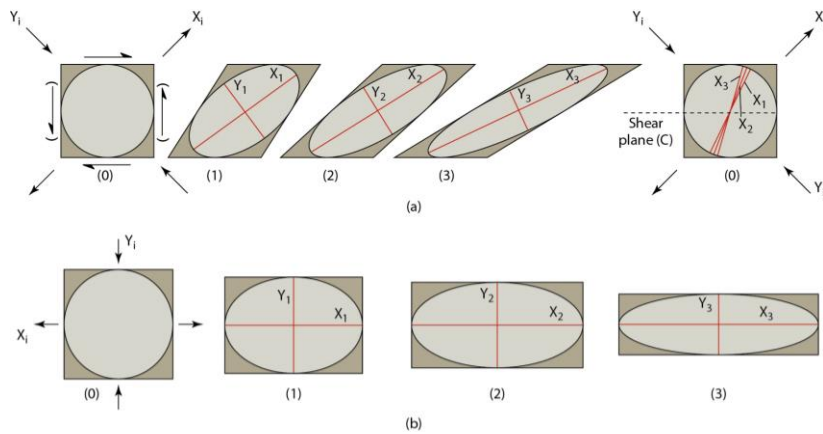
- Homogeneous strain: transformation of square to rectangle, circle to ellipse ("strain ellipse"). Note: lines in circle all equal length; line lengths in square vary by orientation.
- Two material lines perpendicular before and after strain are **principal axes** of strain ellipse (red lines): $X \geq Y \geq Z$.
- Dashed lines are **material lines** that do not remain perpendicular after strain; they rotate toward long axis of strain ellipse (X).

Strain path



- Incremental strain (steps)
- Finite strain (difference between unstrained and final shape)
- Infinitely small strain increments (mathematical): instantaneous strain

Strain Accumulation: coaxial and non-coaxial strain

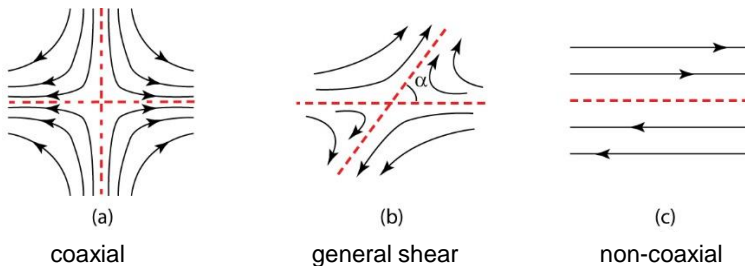


a. (progressive) simple shear or non-coaxial strain

b. (progressive) pure shear, or coaxial strain

Compare incremental strain (X_i, Y_i) and finite strain ($X_{1,2,3}, Y_{1,2,3}$): **vorticity**

Vorticity (internal rotation) and Particle Paths



Rotation of material lines wrt to finite strain is kinematic vorticity.

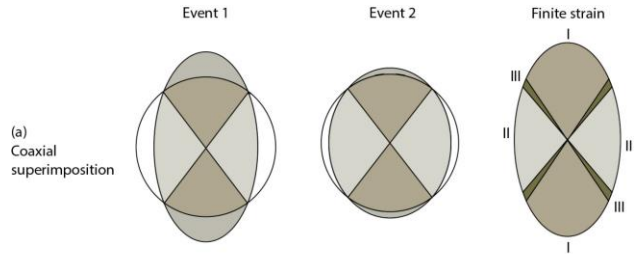
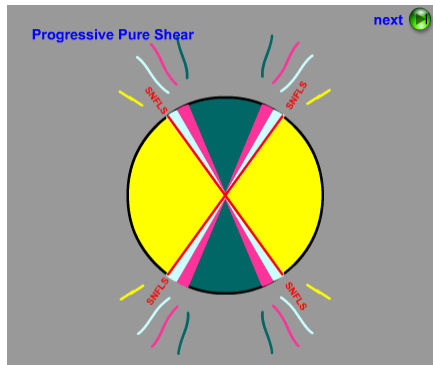
Kinematic vorticity number: $W_k = \cos \alpha$

a. $W_k = 0$ (pure shear, coaxial strain end-member)

b. $0 < W_k < 1$ (general shear, non-coaxial strain)

c. $W_k = 1$ (simple shear, non-coaxial strain end-member)

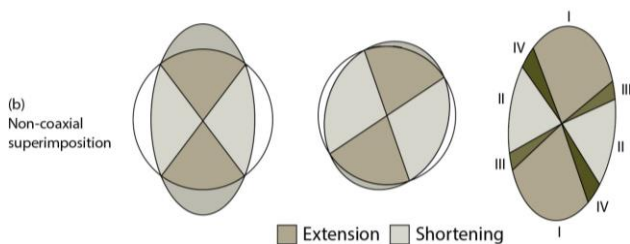
Progressive Coaxial Strain – Pure Shear



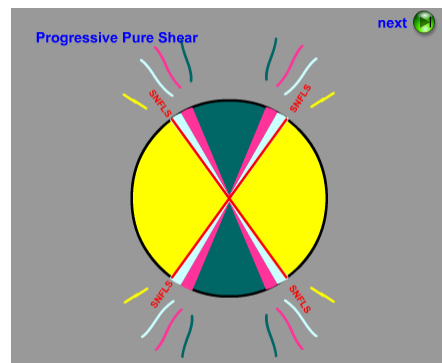
- I. lines continue to be extended
- II. lines continue to be shortened
- III. shortening during event 1 is followed by extension during event 2.

DePaor, 2002

Progressive Non-coaxial Strain – Simple Shear



- I. lines continue to be extended
- II. lines continue to be shortened
- III. shortening during event 1 is followed by extension during event 2.
- IV. extension is followed by shortening during event 2

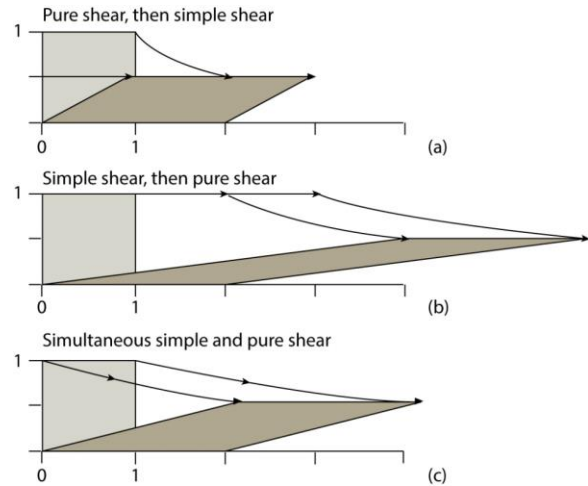


DePaor, 2002

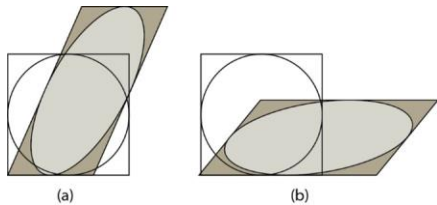
Strain Addition

“Adding” non-coaxial strain and coaxial strain is non-commutative:

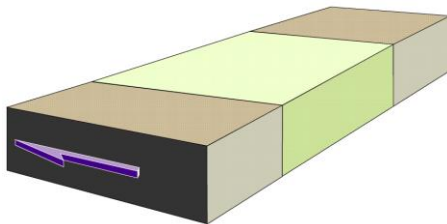
$$\text{strainA} + \text{strainB} \neq \text{strainB} + \text{strainA}$$



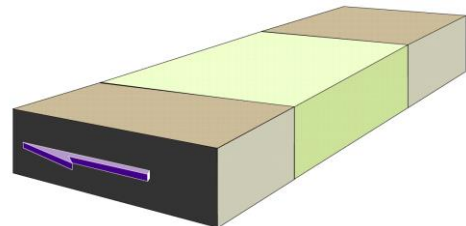
Combinations: Transtension and Transpression



- a. Transtension
- b. Transpression



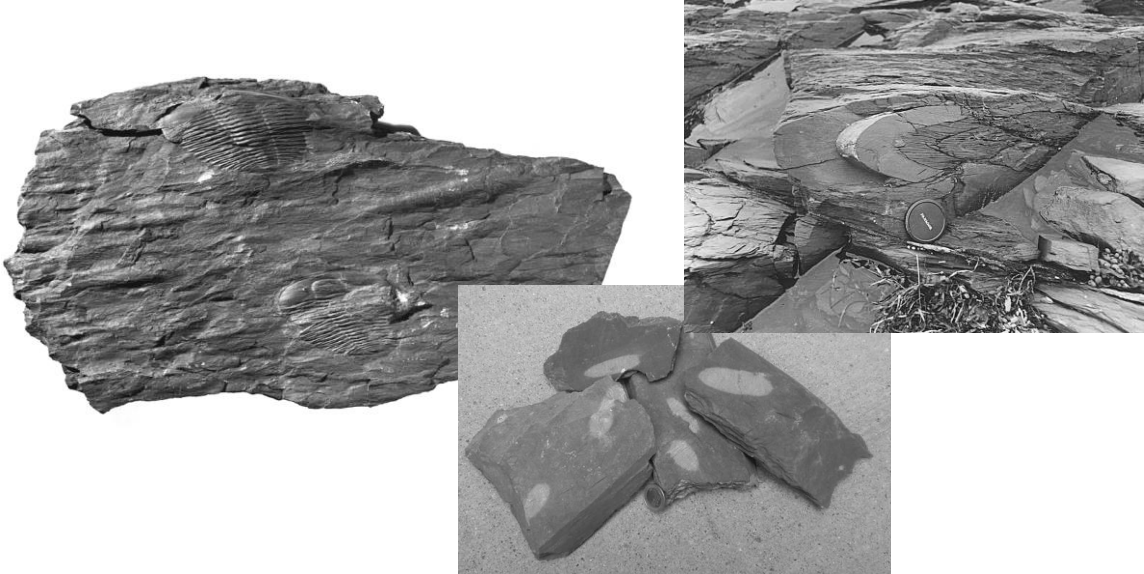
Transtension



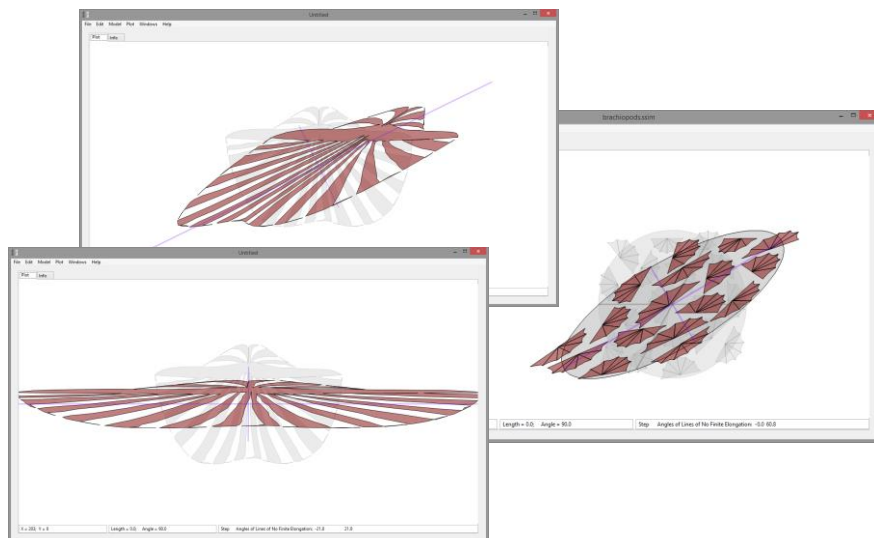
Transpression

Fossen, 2016

Strain Quantification

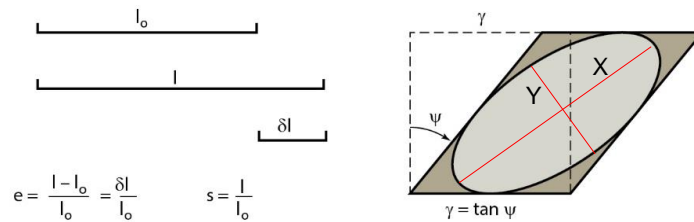


StrainSim (R. Allmendinger)



<http://www.geo.cornell.edu/geology/faculty/RWA/programs/strainsim-v-3.html>

Strain Quantities



e = elongation = $(l - l_0) / l_0 = \delta l / l_0$
 $(e_1 \geq e_2 \geq e_3; e_3 \text{ usually negative})$

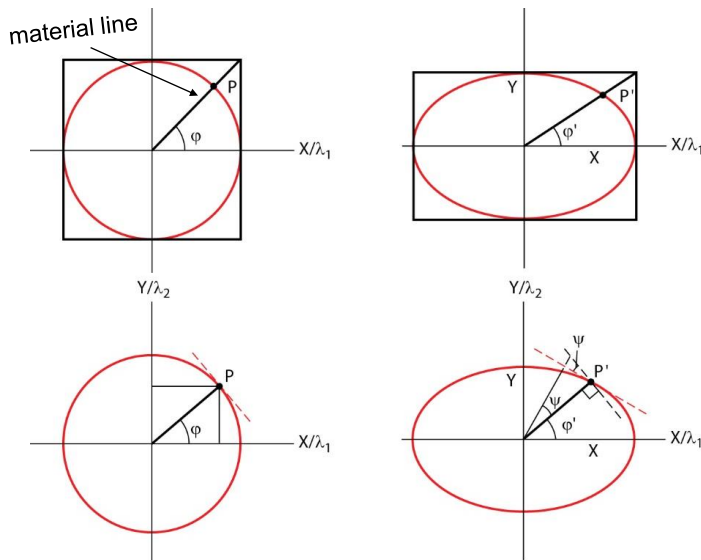
s = stretch = l / l_0 ($X \geq Y \geq Z$; strain ellipsoid axes)
 $= [(l - l_0) / l_0 + l_0 / l_0] = \mathbf{e} + 1$

lambda = quadratic elongation (lambda) = $s^2 = (e + 1)^2$
 $\lambda_1 \geq \lambda_2 \geq \lambda_3$

Misnomer: should be quadratic stretch

γ = shear strain (gamma) = $\tan \psi$
 ψ = angular shear (psi)

Relationship between strain ratio, ψ and line angle and length



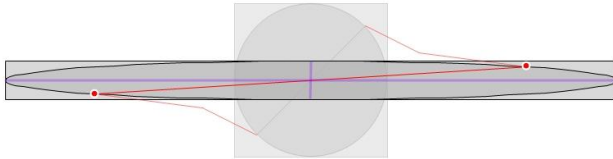
Angle:
 ϕ to ϕ' is \tan of strain ratio

$\tan \phi' / \tan \phi = Y/X$
 $(X \cdot Y = 1, \text{ constant area})$
 So,
 $\tan \phi' = Y/X \cdot \tan \phi$
 $\tan \phi = X/Y \cdot \tan \phi'$

Length:
 l_0 to l is $\sqrt{\text{strain ratio}}$

$l^2 = X^2 + Y^2$
 So,
 $P'(x', y') =$
 $X/Y \cdot \cos^2 \phi, Y/X \cdot \sin^2 \phi$

Strain Calculator



StrainSim

<http://www.geo.cornell.edu/geology/faculty/RWA/programs/strainsim-v-3.html>

Strain Calculator		
strain ratio: X/Y	1	
initial angle, ϕ (deg)	45	45.0
deformed angle, ϕ' (deg)		45.0
initial length, l_0	1	1.0
deformed length, l		1.0

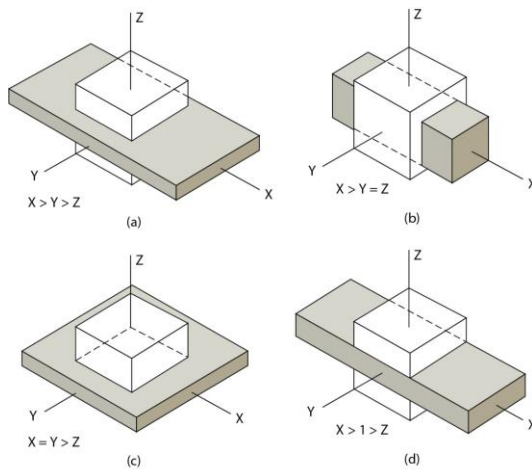
$X=2, Y=1/2$

Strain Calculator		
strain ratio: X/Y	4	
initial angle, ϕ (deg)	45	45.0
deformed angle, ϕ' (deg)		14.0
initial length, l_0	1	1.0
deformed length, l		1.5

$X=4, Y=1/4$

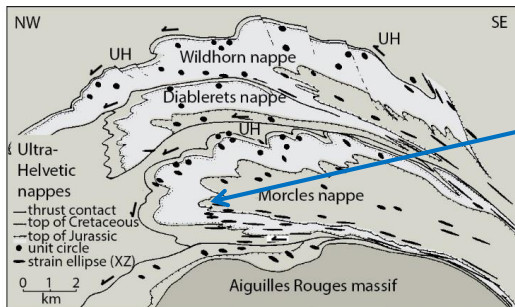
Strain Calculator		
strain ratio: X/Y	16	
initial angle, ϕ (deg)	45	45.0
deformed angle, ϕ' (deg)		3.6
initial length, l_0	1	1.0
deformed length, l		2.8

3D Strain States

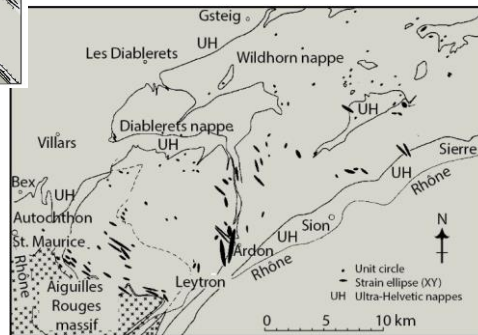


- (a) General strain ($X > Y > Z$)
- (b) Axially symmetric extension ($X > Y = Z$)
- (c) axially symmetric shortening ($X = Y > Z$)
- (d) plane strain ($X > 1 > Z$)
- (e) simple shortening ($1 > Z$)

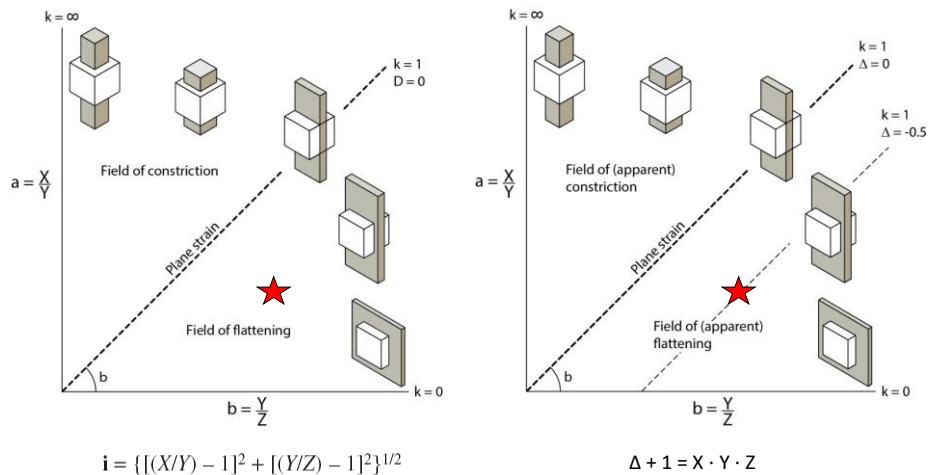
Representation of Strain – Section, Map



Strain in Helvetic “nappes” (large recumbent folds), Switzerland.

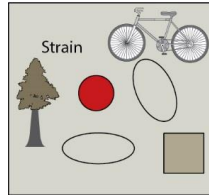


Representation of Strain – 3D in 2D-Space

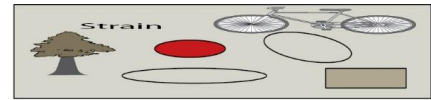


3D strain geometry in 2D plot: axial ratios and volume change

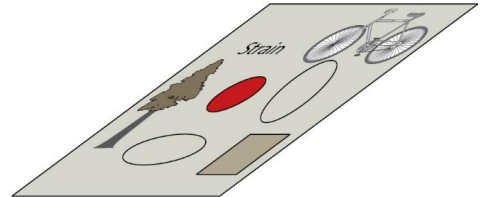
Strain Methods: Variably-shaped Objects



(a)

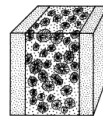


(b)

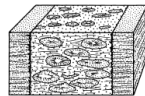


(c)

Strain Methods: Initially Spherical Objects



(a)



(b)

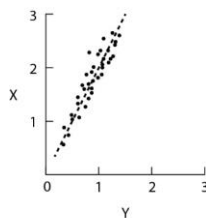


(c)

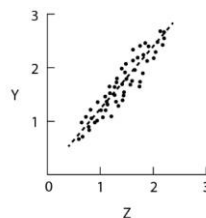
Deformed ooids after:

(b) 25% ($X/Z = 1.8$) shortening

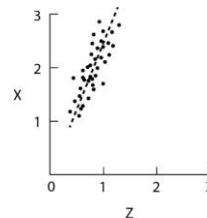
(c) 50% ($X/Z = 4.0$) shortening



$$\frac{X}{Y} = 2$$

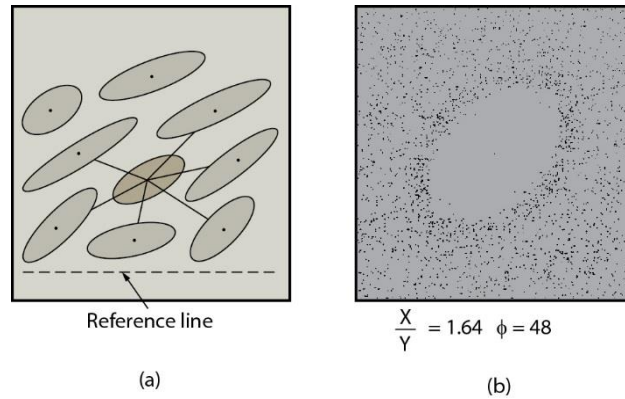


$$\frac{Y}{Z} = 1.2$$



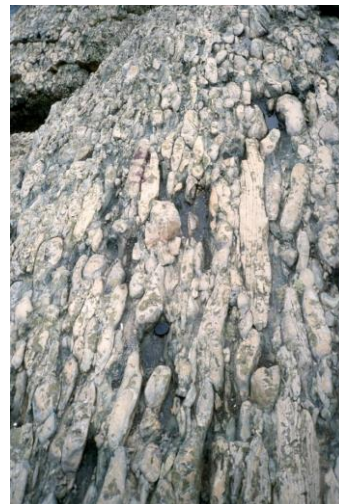
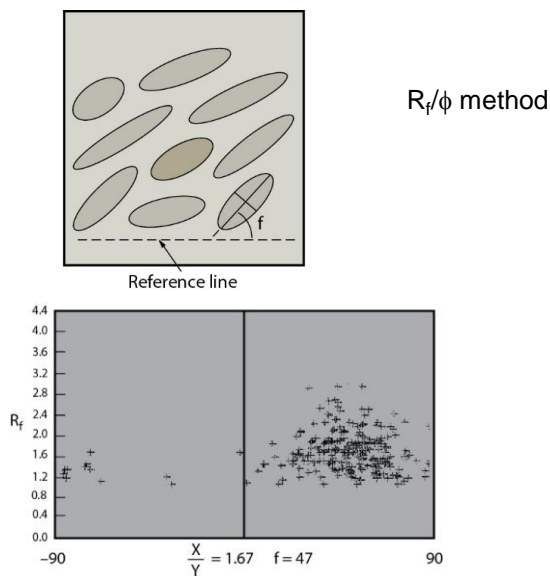
$$\frac{X}{Z} = 2.4$$

Strain Methods: Non-spherical Objects, c-t-c



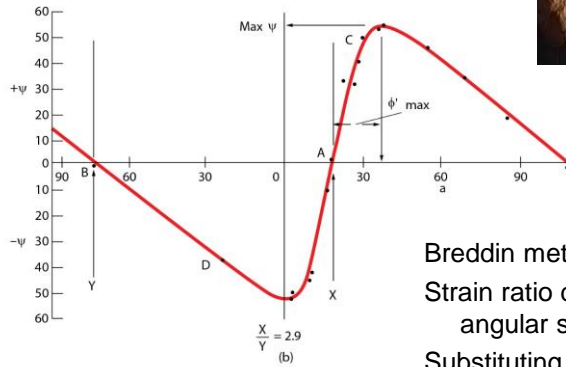
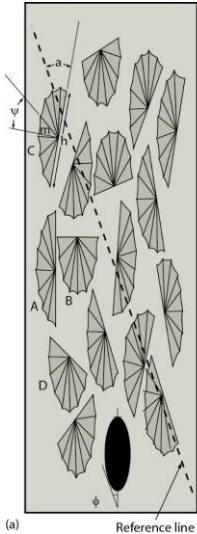
Center-to-center method (or Fry method):
spacing varies as function of finite strain

Strain Methods: Non-spherical Objects, R_f/ϕ



R_f/ϕ method

Strain Methods: Angular Changes



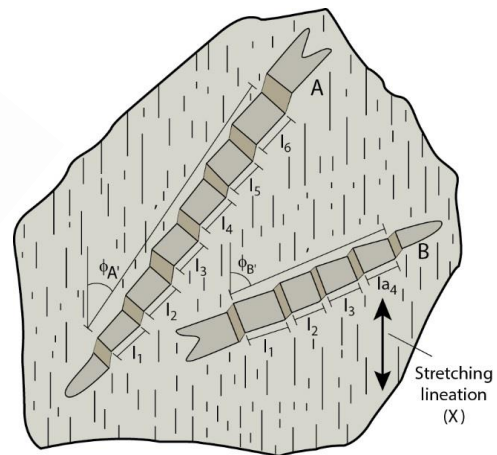
Bredden method:

Strain ratio determined at maximum angular shear, angle $\phi = 45^\circ$

Substituting $\tan \phi = 1$

$$Y/X = \tan \phi'_{\text{max}}$$

Strain Methods: Length Changes



$$\lambda'_A = [(l_1 + l_2 \dots l_6)]^2$$

$$\lambda'_B = [(l_1 + l_2 \dots l_4)]^2$$

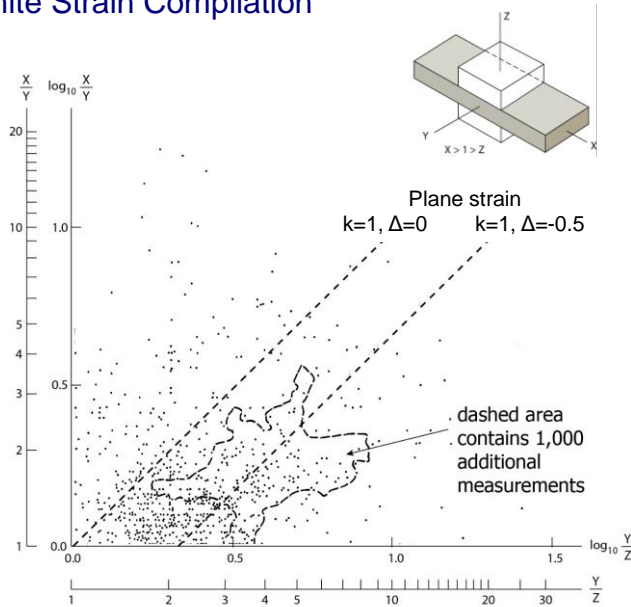
$$X = 1.36$$

$$Y = 1.11$$

$$\Delta = 0.51$$

$$X/Y = 1.23$$

Finite Strain Compilation



Typical strain values:

$$1 < X/Z < 20$$

$$1 < X < 3$$

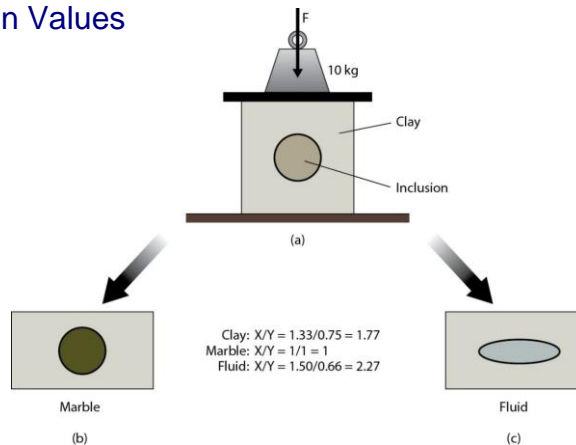
$$0.13 < Z < 1$$

Plane strain: $X > 1 > Z$ ($k=1$)

Δ is volume change

Where does lost rock volume go? Veins.

Understanding Strain Values



Strain and Mechanical Contrast:

- Passive markers have no mechanical contrast: bulk rock strain (clay and clay inclusion; oolite)
- Active markers have mechanical contrast: marker strain (clay and marble or fluid; conglomerate)